

IL-269-10

SITE HYDROGEOLOGY

Stauffer Chemical Company
Chicago Heights, Illinois

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EPA Region 5 Records Ctr.



360895

REFERENCE 4

SITE NAME TRIEM STEEL

SITE ID ILD001744572

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INTRODUCTION

Bruce S. Yare and Associates, Inc. were retained by Stauffer Chemical Company to define hydrogeologic conditions around a closed disposal area at the Chicago Heights, Illinois plant (Figure 1).

From January 8 to 16, 1981, four monitoring wells were installed to increase the areal extent of the existing monitoring well network. The five-inch diameter wells were constructed by drilling 25 to 30 feet to bedrock with a hydraulic rotary rig. Drilling mud was made with bentonite and municipal water. A four-foot long, slotted PVC well screen was attached to steel casing, set in the bottom of the borehole and backfilled with pea gravel. The remaining annulus was filled with neat cement grout. Well development was limited to a short period of air agitation with the rig's compressor.

Information from the monitoring well network, plant production wells, published literature and state records was used to define the water-bearing units and ground-water flow system at the site.

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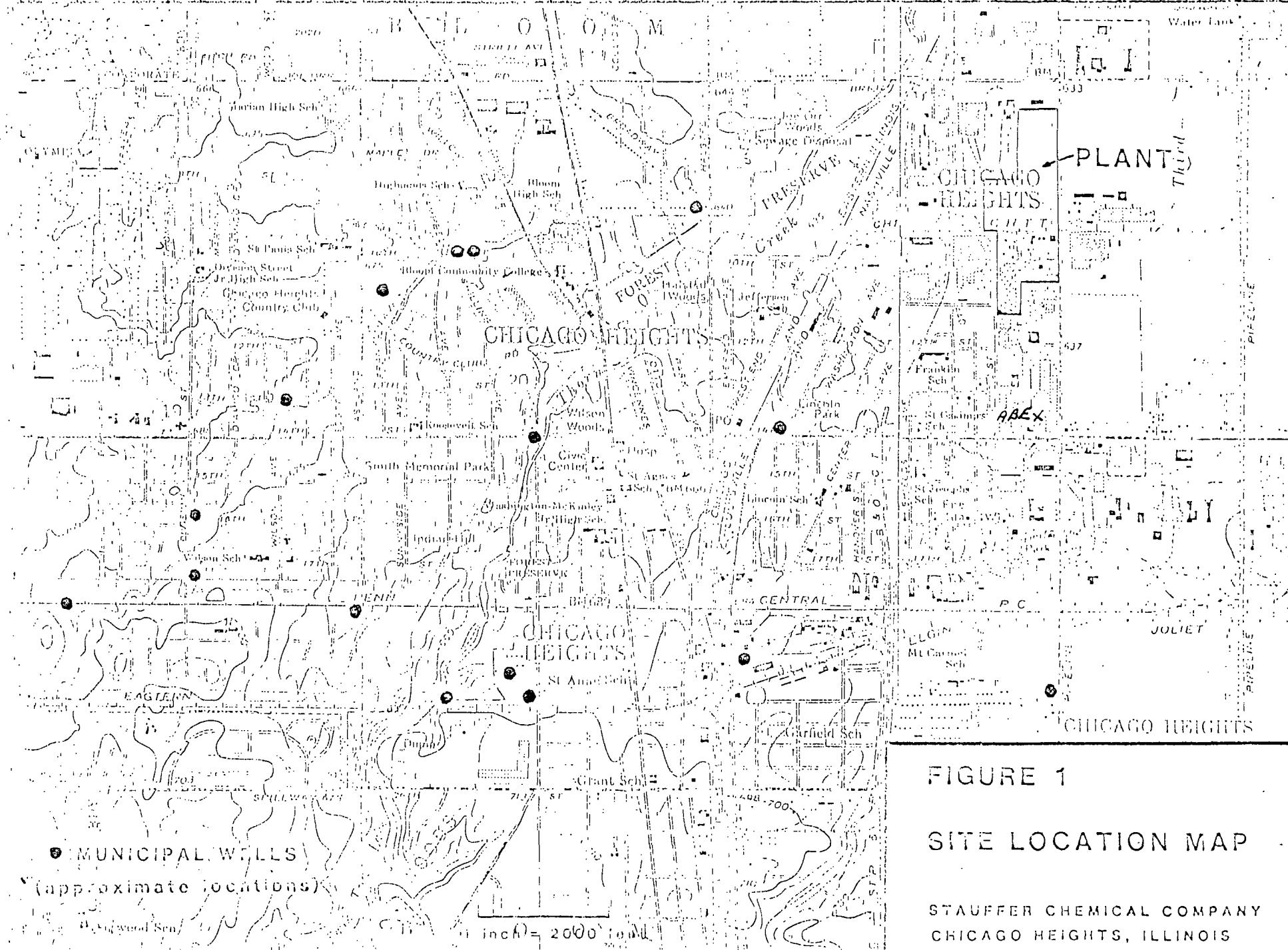


FIGURE 1

SITE LOCATION MAP

STAUFFER CHEMICAL COMPANY
CHICAGO HEIGHTS, ILLINOIS

GROUND-WATER OCCURENCEGlacial Drift

All of northern Illinois is blanketed with unconsolidated sediments deposited by or in front of the numerous ice sheets which once covered the region. These sediments, collectively called glacial drift, consist of till, outwash and glacial lake clays. Tills and glacial lake clays are fine-grained, low permeability sediments. Outwash deposits, composed of sand and gravel, are often important aquifers (Table 1).

Surficial sediments in the vicinity of the plant consist of ground moraine and lake plain (Willman, 1971). Ground moraine is till deposited beneath an ice sheet and a lake plain is a till surface eroded flat by glacial lake wave action.

Regionally-extensive outwash deposits underlie the low permeability till at the site. This sand and gravel deposit is forty feet thick southwest of Chicago Heights and thins toward the northeast (Prickett and others, 1964). Soil borings indicate these deposits range from less than three to more than twenty feet in thickness beneath the plant (Appendix A.1).

The entire drift sequence recharges the underlying dolomite aquifer at a rate of approximately 300,000 gpd/sq mi (gallons per

Table 1. Generalized Stratigraphic and Hydrogeologic Section,
Stauffer Chemical Company, Chicago Heights, Illinois.
(From Hughes and others, 1966.)

SYSTEM	SERIES	MEGA-GROUP	GROUP OR FORMATION	GRAPHIC LOG	THICKNESS (FEET)	DESCRIPTION	AQUIFER SYSTEMS
QUATERNARY	PLEISTOCENE				0 - 400+	Unconsolidated loess and water-sediment deposits; probably clay (ill), silt, sand and gravel, generally discontinuous and interbedded; alluvial soils and sands commonly present along streams.	Glacial drift aquifer system
TRIASSIC					0 - 175	Shale; sandstone, fine-grained; dolomite; sand; clay.	Fractured beds yield small supplies locally.
MESOZOIC					0 - 100+	Dolomite, very pure to very silty, cherry; shale; pebbles; thin shales and siltstones; lenses frequently present in lower parts of Silurian dolomite.	Not consistent; some wells yield more than 100 gpm. Caves and solution channels more abundant near bedrock surface.
DEVONIAN	NIAGARAN				0 - 165		
	ALLEGHENIAN				0 - 250+	Dolomite, sandy at base; light gray to green, plastic to brittle, some dolomitic, silty; dolomite, mostly silty, argillaceous; minor limestone.	
	CHICORINIAN		Mappeets		150 - 350+	Limestone-shale, dark gray, black, brown, plastic to brittle; some dolomite in upper parts; silty, argillaceous.	Yields water from fractured beds. Shales, particularly in lower parts, act as confining beds at the base of the shallow bedrock aquifer system.
ORDOVICIAN	CHAMPLAINIAN	W. 1/2 D.	Celons Plattsburgh		75 - 650	Dolomite, cherry; sandy at base; limestone; shale pebbles.	Where below shales, development and yields of examples are usually where developed by shales, dolomitic sand is fairly common.
						Dolomite, fine to coarse-grained; shale at top; locally cherry and silty at base.	Cherry-colored, fine to hand-powdery, low-intermediate porosity aquifer. Yield about 15% of that of Chambly or Chazy-Chester systems.
	CANADIAN	E. 1/2 D.	Preston du Chien		0 - 340	Dolomite, sandy, cherry, interbedded with sandstone.	The area has dolomite and sandstone generally yielding small amounts of water; a few dolomite locally well-cemented and partially recrystallized may have relatively high yields of several deep wells. Yield about 15% of that of Chambly-Chazy system.
			E. 1/2 D.		0 - 275	Dolomite, white, fine-grained, sandy at base; drusy quartz.	Yield about 15% of Chambly-Chazy system.
			Frankfort		45 - 175	Sandstone, dolomite, and shale; glauconitic; green to red; argillaceous.	
			Gatesville		100 - 275	Sandstone, fine to medium-grained, well sorted, upper part dolomitic.	Yield about 15% of Chambly-Chazy system.
CAMBRIAN	CROIXAN				235 - 450	Shale and siltstone; dolomite; glauconitic; sandstone; dolomite; glauconitic; dolomite, sandy, micaceous.	Shales generally not water-bearing, act as confining bed at the base of the Canadian-Ordivician aquifer system.
			Eau Claire				
			Mt. Simon		2000+	Sandstone, coarse-grained, white, red in lower half; lenses of shale and dolomite, red, micaceous.	Mt. Simon aquifer system
PRECAMBRIAN						Not penetrated by wells in Chazy area. Highly weathered and/or gray dolomite or similar rocks.	Mt. Simon conditions. Yield about probably less permeable than pre-Cambrian; quality of water deteriorates with depth.

day per square mile). However, the hydraulic connection between the two units appears limited because water levels in the glacial drift do not usually coincide with the potentiometric surface in the dolomite aquifer (Prickett and others, 1964).

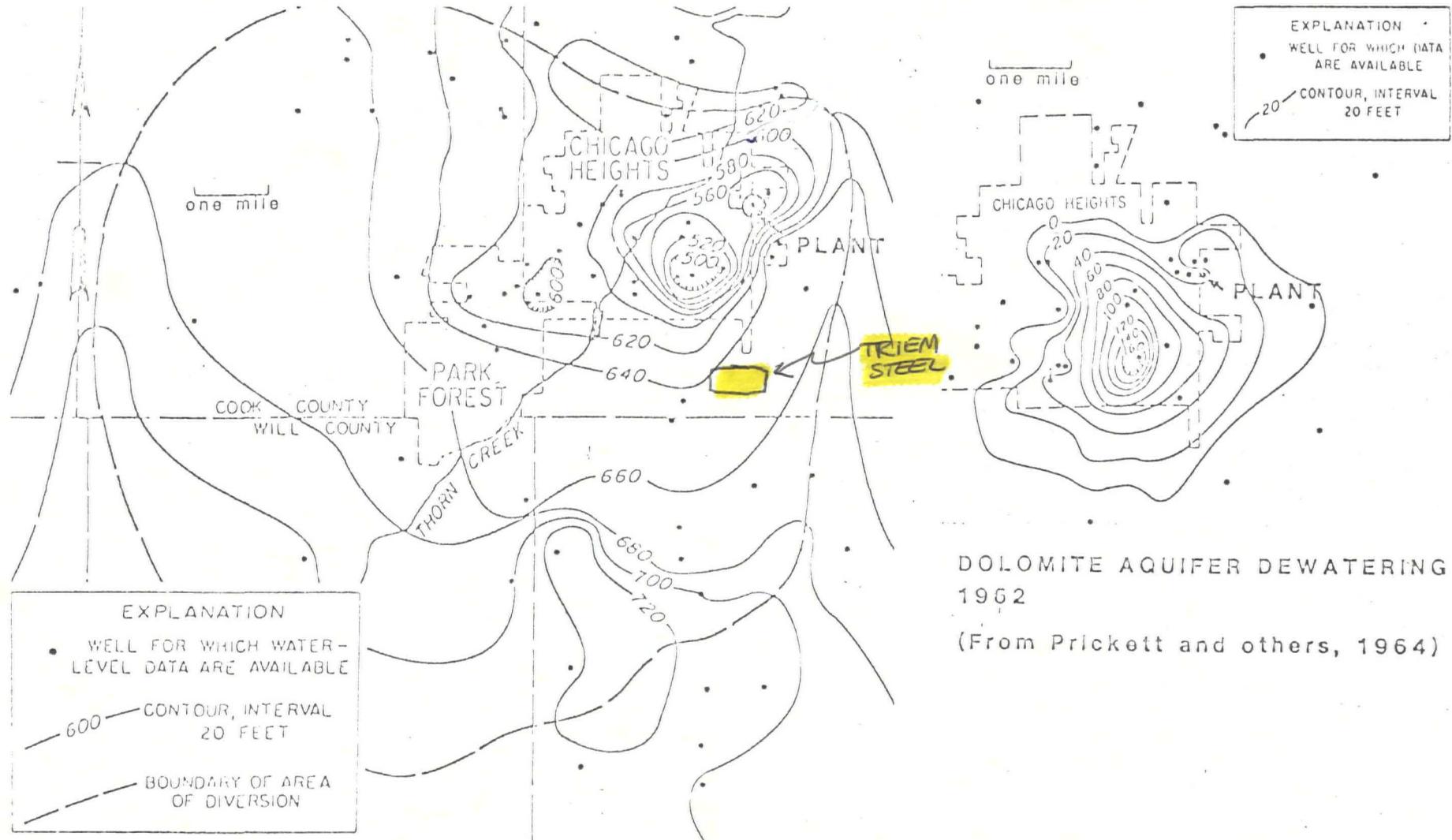
Silurian Dolomite Aquifer

A Silurian-age dolomite formation occurs beneath the drift deposits covering the site (Table 1). This dolomite, approximately 400 feet thick, is a regionally-important aquifer tapped by many high-capacity municipal and industrial wells (Figure 1). Individual well yields of over 1000 gpm (gallons per minute) are not uncommon. Ground water occurs mostly under confined, leaky artesian conditions but there are some areas of water table (unconfined) conditions. In the Chicago Heights area, leaky artesian conditions prevail except where heavy pumping has dewatered the upper portion of the aquifer (Figure 2).

Water levels in the dolomite aquifer show a rapid response to recharge from precipitation. However, surficial sediments and/or low permeability beds in the upper part of the dolomite aquifer probably retard downward movement of water toward the main producing zones. Estimates of the aquifer recharge rate range from 177,000 to 225,000 gpd/sq mi (McDonald and Sasman, 1966; Prickett and others, 1964).

Average aquifer transmissivity at the plant is 21,700 gpd/ft (gallons per day per foot) and ranges from 2,000 to 41,000 gpd/ft

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WATER LEVEL ELEVATION IN THE SILURIAN DOLOMITE AQUIFER, 1966. (From McDonald and Sasman, 1966)

DOLOMITE AQUIFER DEWATERING,
1962

(From Prickett and others, 1964)

FIGURE 2.
REGIONAL HYDROGEOLOGIC
CONDITIONS

STAUFFER CHEMICAL COMPANY
CHICAGO HEIGHTS, ILLINOIS

(Table 2). The long-term gravity yield of the aquifer is 0.03.

In the past, the Chicago Heights plant relied mainly on the shallow dolomite aquifer to meet its water supply needs (Appendix A.2). Withdrawals from the plant production wells had a noticeable impact on the regional ground-water flow system. In 1966, this pumpage produced a large cone of depression at the plant (Figure 2). The regional flow system was controlled by withdrawals from the Inland Steel and Chicago Heights Municipal well fields located southwest of the plant.

Currently, PW-10 is the only plant production well using the dolomite aquifer as a water supply source. Recent water-level maps for the dolomite aquifer are not available, so the present effects of withdrawals from PW-10 are not known.

Cambro-Ordovician Aquifer

The Cambro-Ordovician aquifer is not a massive water-bearing unit but rather consists of a series of interbedded low-yield dolomite formations and high-yield sandstones. Regionally, the aquifer is very heavily developed with large-scale water supplies obtained from the Ironton-Galesville Sandstone and the Glenwood-St. Peter Sandstone (Table 1). It is not extensively used in the Chicago Heights area (Sasman, 1977), and PW-2 is the only plant production well tapping the Cambro-Ordovician aquifer (Appendix A.3).

Aquifer transmissivity in a municipal well near the plant is

Table 2. Silurian Dolomite Aquifer Hydraulic Characteristics,
Stauffer Chemical Company, Chicago Heights, Illinois.
(Data from Prickett and others, 1964.)

<u>Well</u>	<u>Date of Test</u>	<u>Static Water Level</u> (feet)	<u>Discharge Rate</u> (gpm)	<u>Adjusted Specific Capacity</u> (gpm/ft)	<u>Transmissivity</u> (gpd/ft)
PW-1	1921	35	340	5.90	12,000
	1946	36	350	4.50	9,000
PW-3	1941	37	130	1.05	2,000
PW-4	1947	52	350	11.50	23,000
PW-5	1956	49	520	17.35	34,700
PW-6	1955	102	70	5.38	11,000
PW-7	1955	110	130	2.17	—
PW-8	1955	97	250	20.80	41,000
PW-9	No Data	-	-	-	—
PW-10	1956	95	500	20.50	41,000

Note: PW-2 which taps the Cambro-Ordovician aquifer was tested in 1942. At a discharge of 680 gallons per minute the specific capacity was reported as 5.3 gpm/ft.

1,600 gpd/ft. The average regional storage coefficient is 0.00035 for short time periods and 0.0006 for periods of several years or more (Suter and others, 1959).

A thick confining bed, the Maquoketa Shale, separates the Cambro-Ordovician aquifer from the overlying Silurian dolomite aquifer. The low vertical permeability of the shale ($0.00005 \text{ gpd}/\text{ft}^2$) limits but does not prevent the downward movement of water under natural and pumping-induced gradients. Under a head differential of several hundred feet, the Cambro-Ordovician aquifer receives 1500 gpd/sq mi of recharge from the overlying dolomite aquifer (Prickett and others, 1964).

GROUND-WATER FLOW SYSTEM

Water levels in the plant monitoring and production wells were measured several times during May and August, 1981 (Appendix B). A water-level elevation map was compiled using measurements made from May 18 to 21, 1981 (Figure 3). On the basis of this map, the ground-water flow direction in the glacial drift aquifer is toward the south and southwest.

The hydraulic gradient is not uniform across the site, changing from 0.00048 north of the disposal area to 0.0076 south of it. Assuming a horizontal permeability of 25 ft/day (feet per day) for fine, silty sands and a porosity of 35 percent, the ground-water flow rate in the glacial drift aquifer is:

$$V = \frac{KI}{n}$$

where: V = velocity, ft/day
K = permeability, ft/day
I = ground-water gradient
n = formation porosity

North Area

$$V = \frac{25 \text{ ft/day}(0.0048)}{0.35}$$

$$V = 0.3 \text{ ft/day}$$
$$= 110 \text{ ft/year}$$

South Area

$$V = \frac{25 \text{ ft/day}(0.0076)}{0.35}$$

$$V = 0.5 \text{ ft/day}$$
$$= 182 \text{ ft/year}$$

The observed change in ground-water gradient is probably due to the influence of withdrawals from PW-10 or some other nearby

pumping center. Pumping from production well PW-2 should not influence water levels in the glacial drift because these sediments, as well as the Silurian dolomite aquifer, are cased off.

Head differences in wells tapping the glacial drift aquifer and the Silurian dolomite aquifer indicate the hydraulic connection between the two water-bearing units may be limited. Water levels in production wells PW-3 and PW-9 are one to three feet lower than water levels in nearby monitoring wells (Figure 3).

Other evidence of a limited hydraulic connection is the apparent lack of drawdown in the drift aquifer due to pumpage from the dolomite aquifer. In May 1981, production well PW-10 was pumping at a rate of 118 gpm and significant drawdowns were observed in dolomite aquifer wells located 250 to 400 feet away. A comparable drawdown was not observed in monitoring well MW-11 tapping the drift aquifer 400 feet southeast of PW-10. (Figure 3).

From the information available, it appears that: 1) withdrawals from PW-10 do not significantly effect water levels and ground-water flow patterns in the glacial drift aquifer, 2) withdrawals from production well PW-2 do not significantly effect water levels in the Silurian dolomite aquifer and 3) flow in the glacial drift aquifer is not controlled by pumpage from the plant's wells.

Respectfully Submitted,

Bruce S. Yare and Associates, Inc.

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President

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Appendix A.1. Representative Soil Borings, Stauffer Chemical Company, Chicago Heights, Illinois. (Borings done in 1976 and 1977 by Walter H. Flood & Co., Hillside, Illinois. All measurements in feet below grade.)

BORING No. 2 (Location - 10+50N, 12+68E)

<u>Depth</u>	<u>Description</u>
0.0 - 0.5	GRAVEL
0.5 - 4.5	CONCRETE
4.5 - 15.0	CLAY - brown and gray, silty
15.0 - 25.0	CLAY and SILT - gray
25.0 - 28.5	SAND - fine to medium, with small to large gravel
28.5 - 38.5	BEDROCK - weathered dolomite

BORING No. 6 (Location - 12+65N, 12+60E)

<u>Depth</u>	<u>Description</u>
0.0 - 6.5	FILL
6.5 - 11.0	SILT - brown and gray, clayey
11.0 - 30.0	SAND - fine to medium, brown
30.0 - 33.0	BEDROCK - weathered dolomite

BORING No. 9 (Location - 16+15N, 12+85E)

<u>Depth</u>	<u>Description</u>
0.0 - 6.0	FILL
6.0 - 15.5	SILT - brown and gray, clayey
15.5 - 31.0	SAND - fine, gray
31.0 - 31.5	BEDROCK - gray dolomite

BORING NO. 10 (Location - 16+00N, 11+40E)

<u>Depth</u>	<u>Description</u>
0.0 - 5.0	FILL
5.0 - 6.0	CLAY LOAM - black
6.0 - 20.0	CLAY - brown to gray, silty
20.0 - 23.0	CLAY and SILT - interbedded, gray
23.0 - 30.0	SAND - fine, gray
30.0 - 33.0	SILT - gray, some fine sand
33.0 - 35.0	BEDROCK - gray dolomite

Appendix A.2. Geologic Log of Production Well PW-3, Stauffer Chemical Company, Chicago Heights, Illinois. (Data from Illinois State Geological Survey files. All measurements in feet.)

<u>PLEISTOCENE SYSTEM</u>	<u>Thickness</u>	<u>Depth</u>
DRIFT and FILL	36	36

SILURIAN SYSTEM

Niagaran Series

DOLOMITE	- broken, slightly cherty, greenish gray to buff, very fine	9	45
SILTSTONE	- broken, dolomitic, gray	19	64
DOLOMITE	- creviced, pink to gray, silty	58	122
DOLOMITE	- hard, very silty, greenish gray	30	152
DOLOMITE	- soft, very silty, greenish gray, gray shale streaks	27	179
DOLOMITE	- some crevices, gray to buff, fine, dense	50	229
DOLOMITE	- greenish gray, fine, dense	25	254
DOLOMITE	- white to gray, fine, dense water level dropped 12 feet at a depth of 291 feet	146	400

Alexandrian Series

Kankakee Formation

DOLOMITE	- hard, buff, fine to medium, glauconitic, pyritic, slightly vesicular	33	433
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ORDOVICIAN SYSTEM

Edgewood Formation

SILTSTONE	- dolomitic, gray	—	433
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Appendix A.3. Geologic Log of Production Well PW-2, Stauffer Chemical Company, Chicago Heights, Illinois. (Data from Illinois State Geological Survey files. All measurements in feet.)

<u>PLEISTOCENE SYSTEM</u>		<u>Thickness</u>	<u>Depth</u>
GLACIAL DRIFT(?)		36	36
<u>SILURIAN SYSTEM</u>			
LIME	- blue, medium hard	69	98
LIME	- white, medium hard	16	114
LIME	- blue, medium hard	21	135
LIME	- white	130	265
LIME	- gray, fairly soft	165	430
<u>ORDOVICIAN SYSTEM</u>			
<u>Richmond Formation</u>			
SHALE		240	670
<u>Galena-Platteville Formation</u>			
LIME	- white, hard	335	1005
<u>St. Peter Formation</u>			
SAND		120	1125
CAVE	- no casing needed	25	1150
<u>Oneota Formation</u>			
LIME	- white, hard	150	1300
<u>CAMBRIAN SYSTEM</u>			
<u>Trempealeau Formation</u>			
LIME	- blue gray, very hard	130	1480

Appendix A.3. Geologic Log of Production Well PW-2, Stauffer
Chemical Company, Chicago Heights, Illinois.
(continued)

<u>CAMBRIAN SYSTEM (continued)</u>	<u>Thickness</u>	<u>Depth</u>
<u>Franconia Formation</u>		
SHALE - sandy	105	1585
SHALE - limey, hard	9	1594
SHALE - soft	21	1615
<u>Dresbach Formation</u>		
SAND - with some lime rock	70	1685
SAND - Potsdam	105	1790
<u>Eau Claire Formation</u>		
SAND and LIME	7	1797

Appendix B. Water-Level Information, Stauffer Chemical Company,
 Chicago Heights, Illinois. (Depth to water in feet
 below top of casing and elevation in feet above mean
 sea level.)

<u>WELL</u>	<u>May 18 to 21, 1981</u>		<u>August 5 to 6, 1981</u>	
	<u>Depth to Water</u>	<u>Elevation</u>	<u>Depth to Water</u>	<u>Elevation</u>
MW-1	30.01	615.67	28.68	617.00
MW-2A	28.21	612.45	26.93	613.73
MW-4	25.92	614.14	24.87	615.19
MW-5	23.03	614.53	21.81	615.75
MW-7	16.30	618.78	15.95	619.13
MW-8	10.56	623.43	11.47	622.52
MW-9	12.25	622.02	12.32	621.95
MW-10	18.73	618.44	17.52	619.65
MW-11	27.54	610.69	26.65	611.58
MW-12	41.28	602.16	39.63	603.81
PW-2 ⁽¹⁾	137	?	139	?
PW-3	32.58	607.39	31.41	608.56
PW-4	45.58	591.49	40.52	596.55
PW-6	36.75	?	34.56	?
PW-7	52.64	586.57	46.71	592.50
PW-8	29.91	?	28.07	?
PW-9	11.94	621.28	12.00	621.22

Note: 1) Measurement is airline reading in feet, length of airline not known. Sasman and others (1977) reported a water level elevation of 133 feet above msl for this well.